

Centre for Hydrology Coldwater Laboratory,
Canmore, Canada



UAV-borne LiDAR Observations at the Canadian Rockies Hydrological Observatory

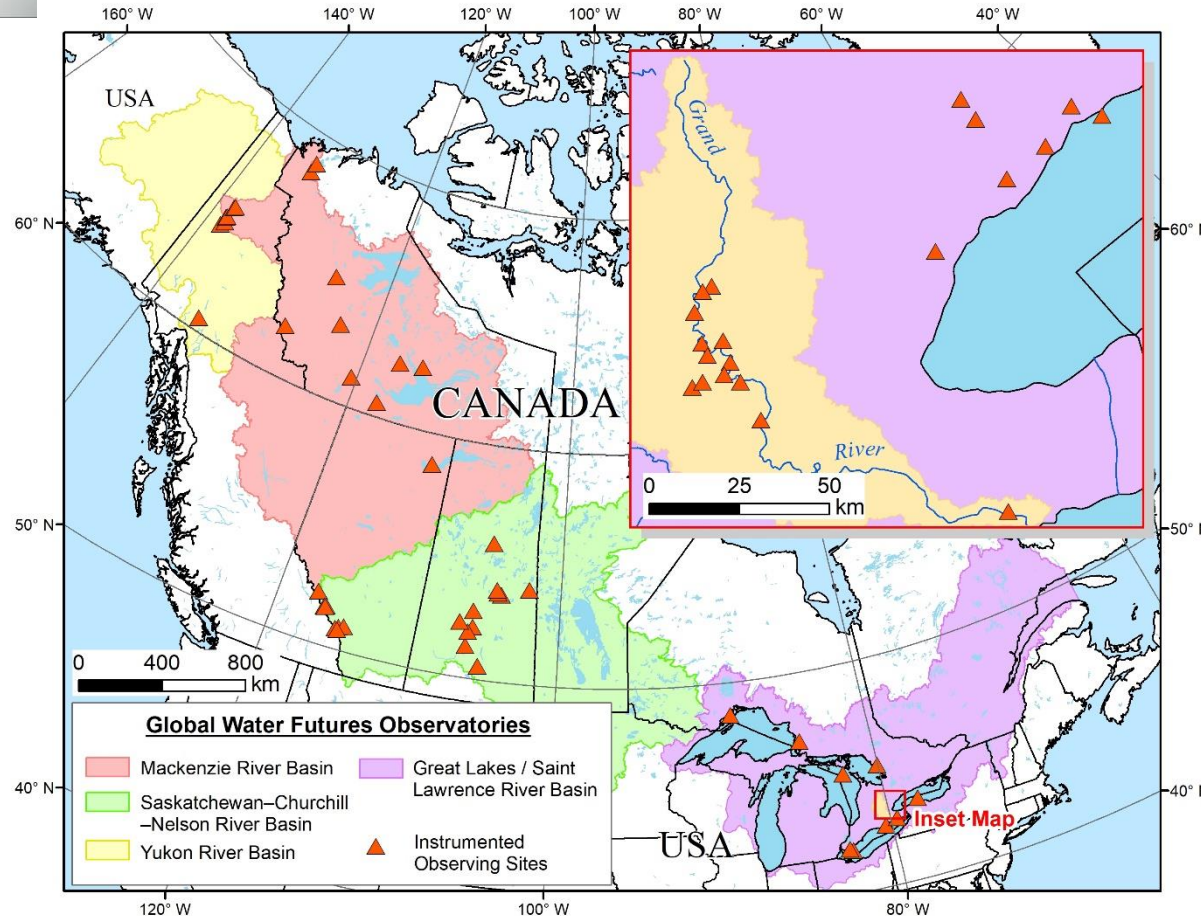
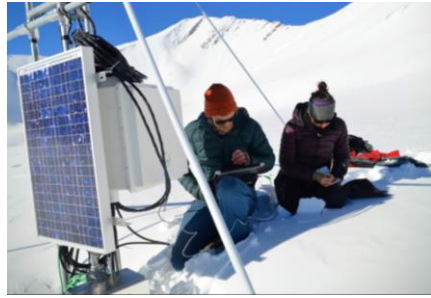
John Pomeroy, Madison Harasyn,
Kieran Lehan, Hannah Koslowsky,
Lindsey Langs, Xing Fang

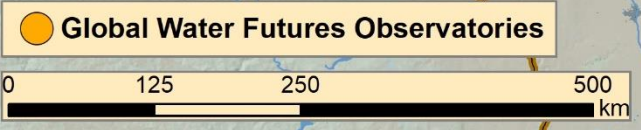
Centre for Hydrology,
University of Saskatchewan,
Canmore, Canada

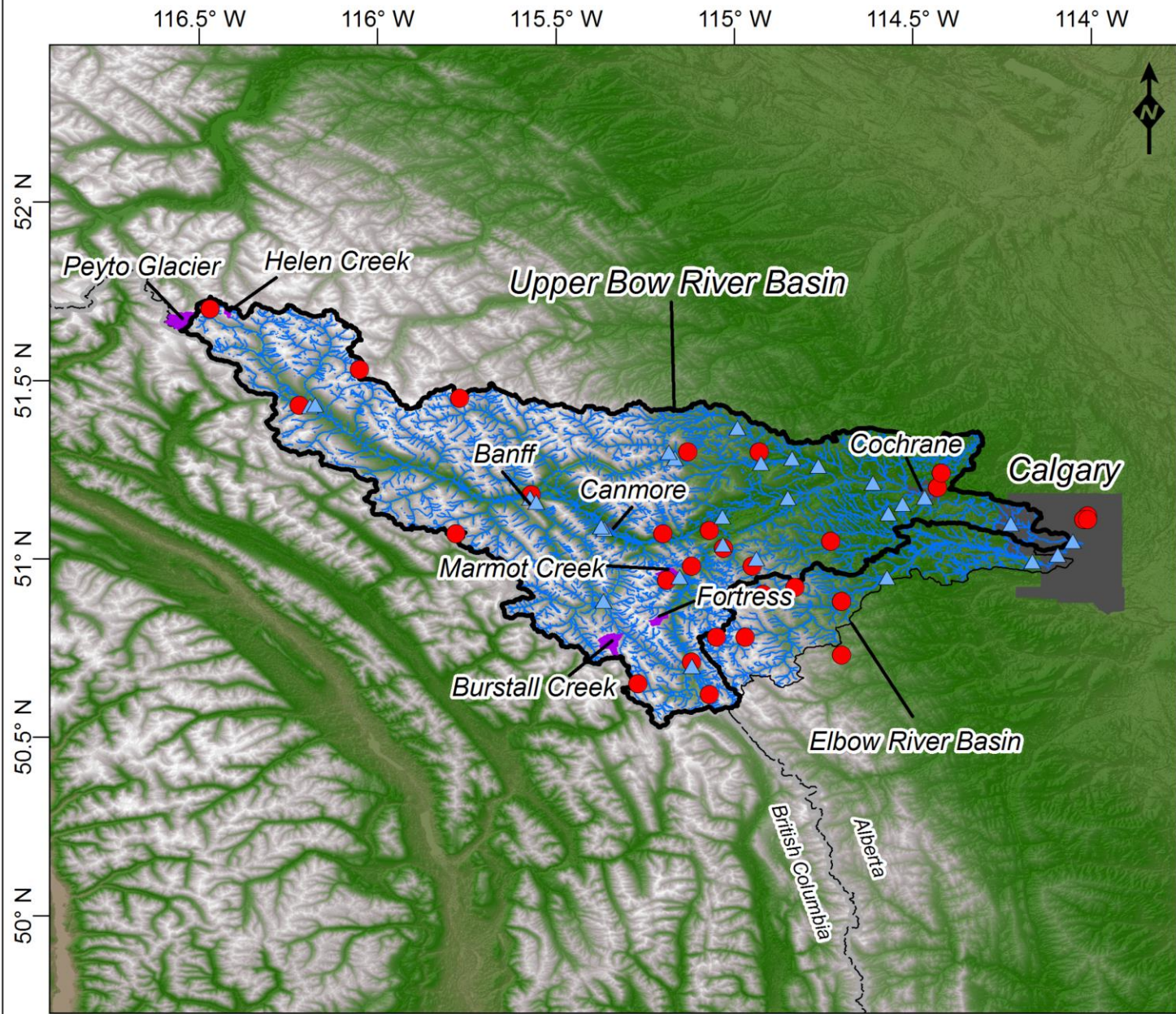











Canada's freshwater early warning system



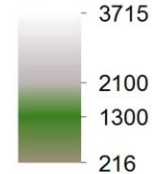




Legend

-  Hydrometric Stations
-  Hydrometeorological Stations
-  Upper Bow River Basin
-  Research Basins
-  Drainage
-  Cities
-  Province Boundaries

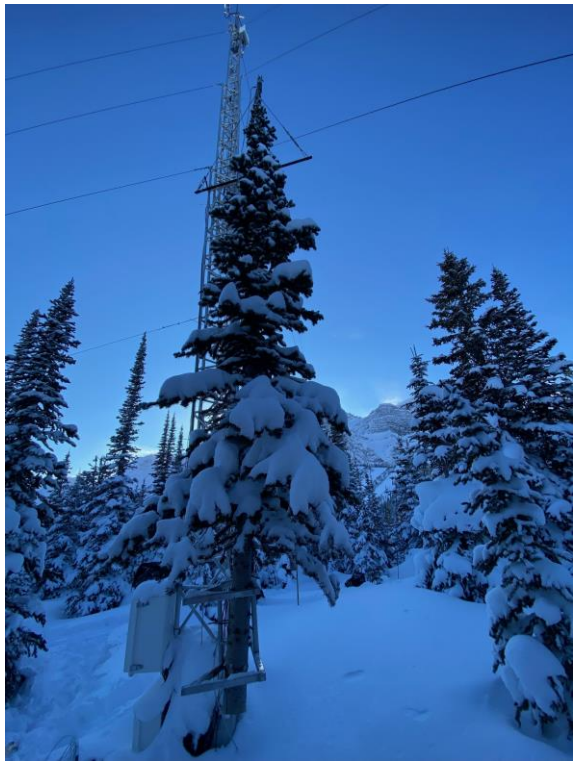
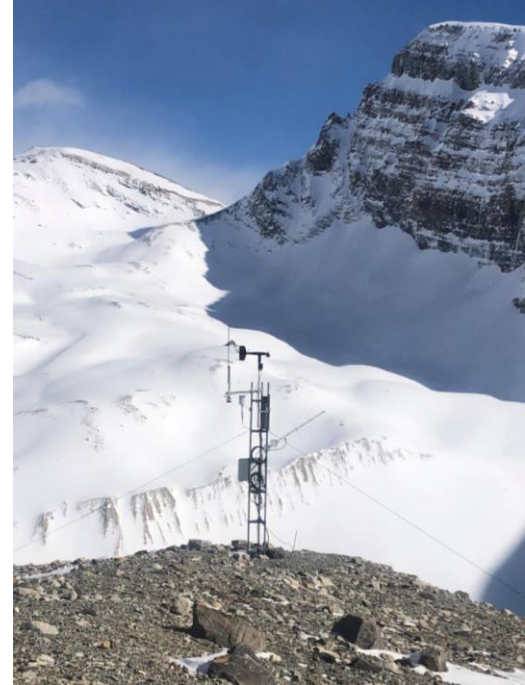
Elevation (m)



Data Source: Canadian Digital Elevation Model and National Hydro Network (Environment Canada)



Canadian Rockies Hydrological Observatory



Glacier Monitoring at Peyto Glacier



Glacier Flow Rate

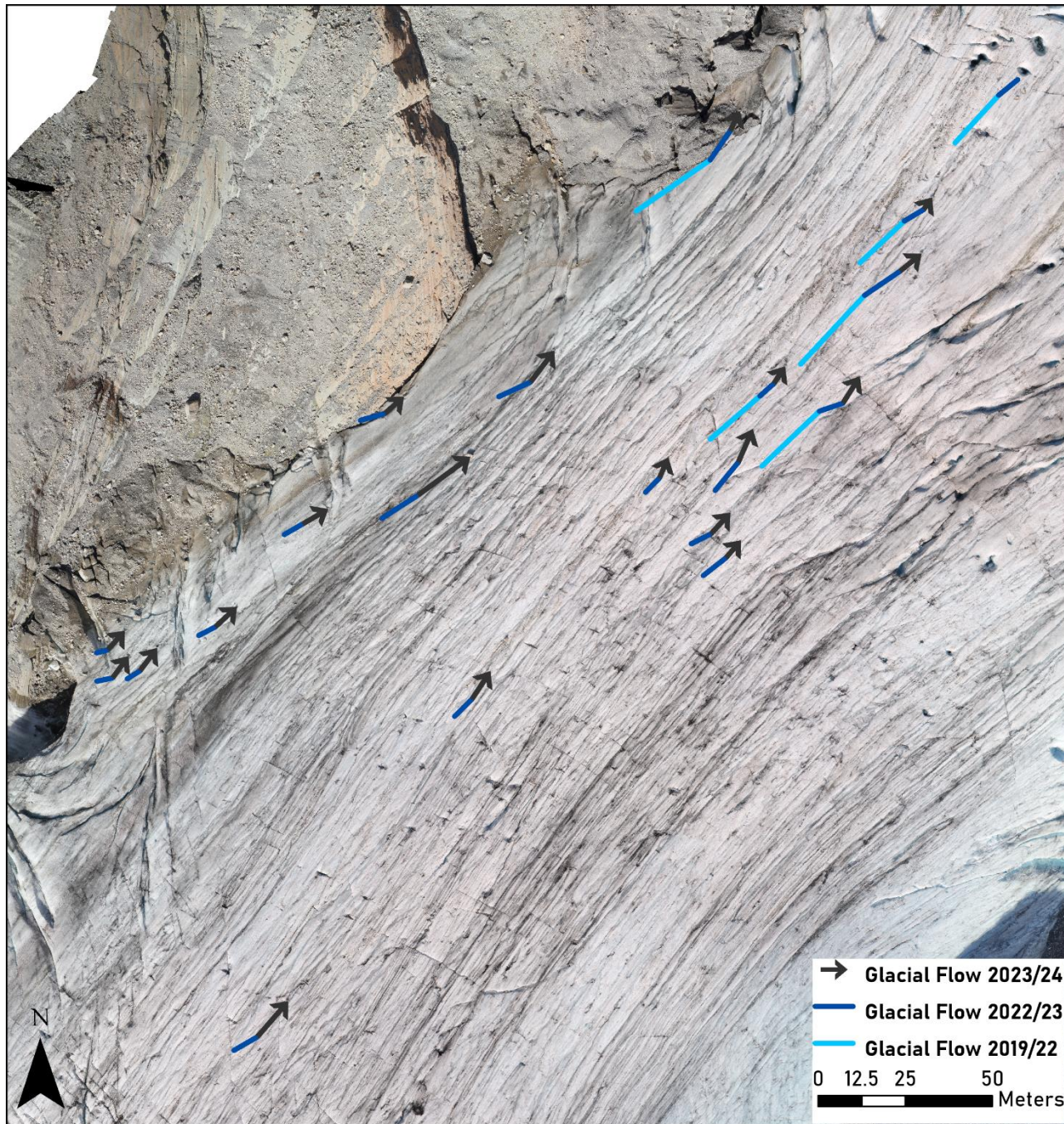
Glacier flow measured by locating rocks on the surface of the upper glacier in RGB photography from different years.

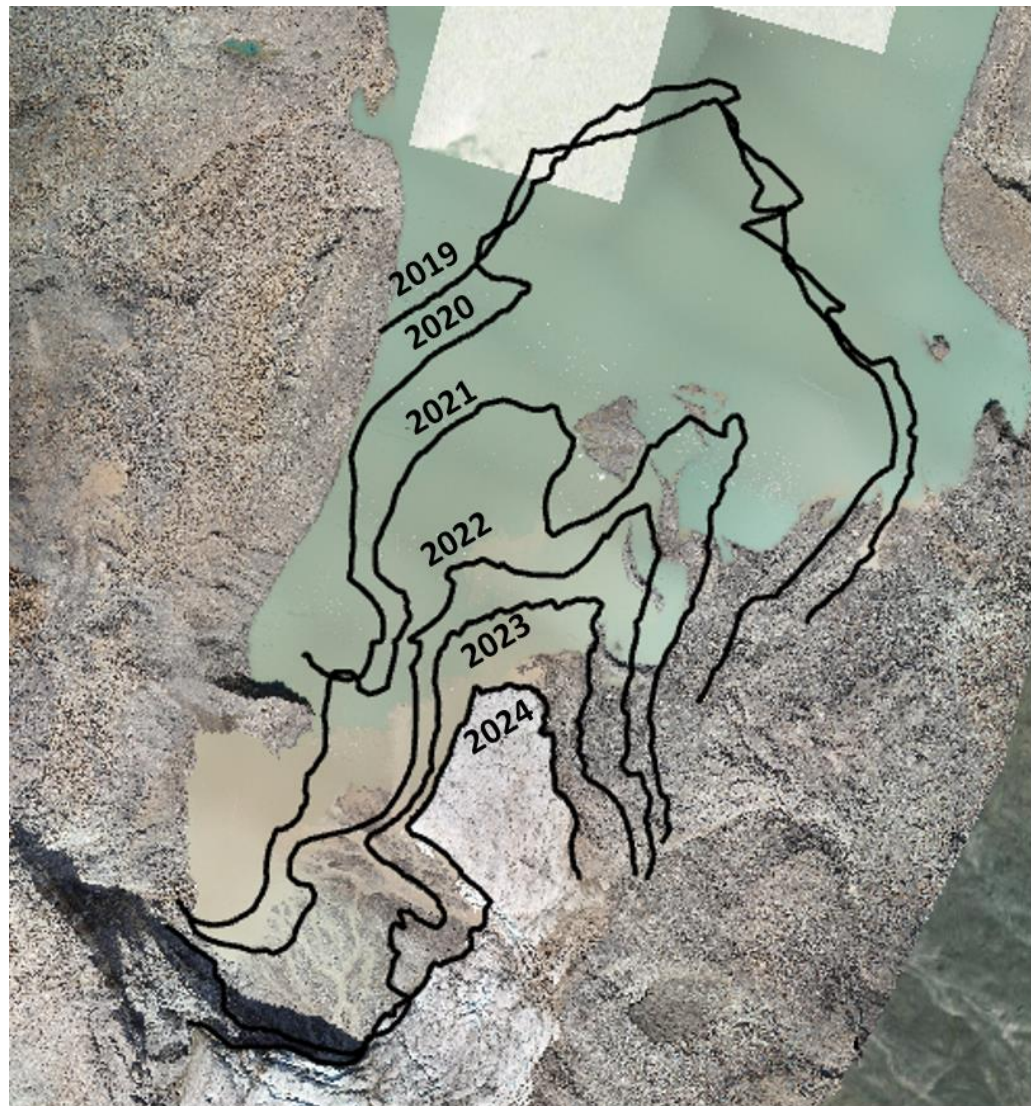
Average flow:

2019-2022: 7.27m/year

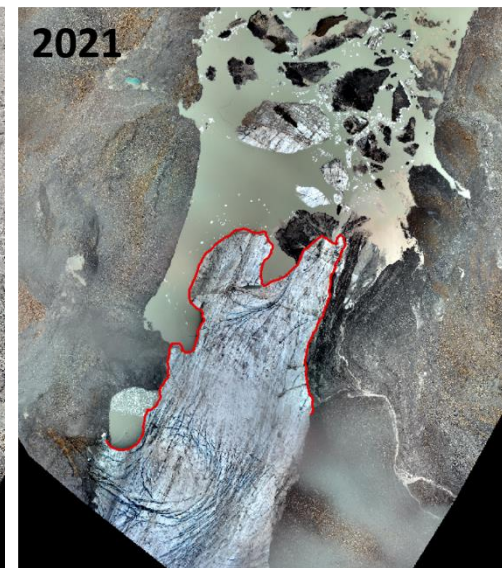
2022-2023: 7.13m/year

2023-2024: 8.45m/year





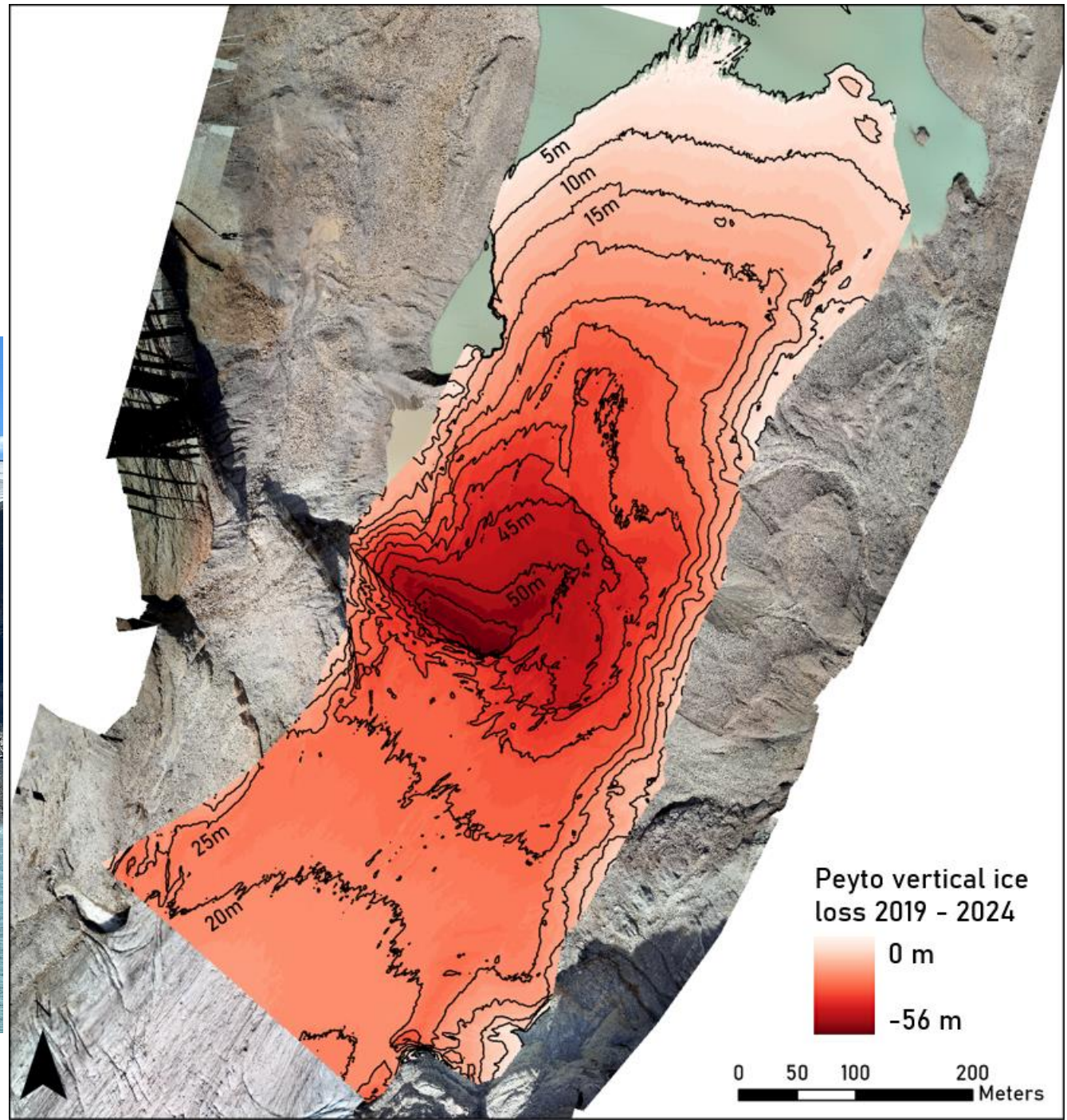
Avg. Lateral Retreat
 2019 – 2020: -12.5 m
 2020 – 2021: -195.7 m
 2021 – 2022: -120.7 m
 2022 – 2023: -80.0 m
 2023 – 2024: -34.37 m



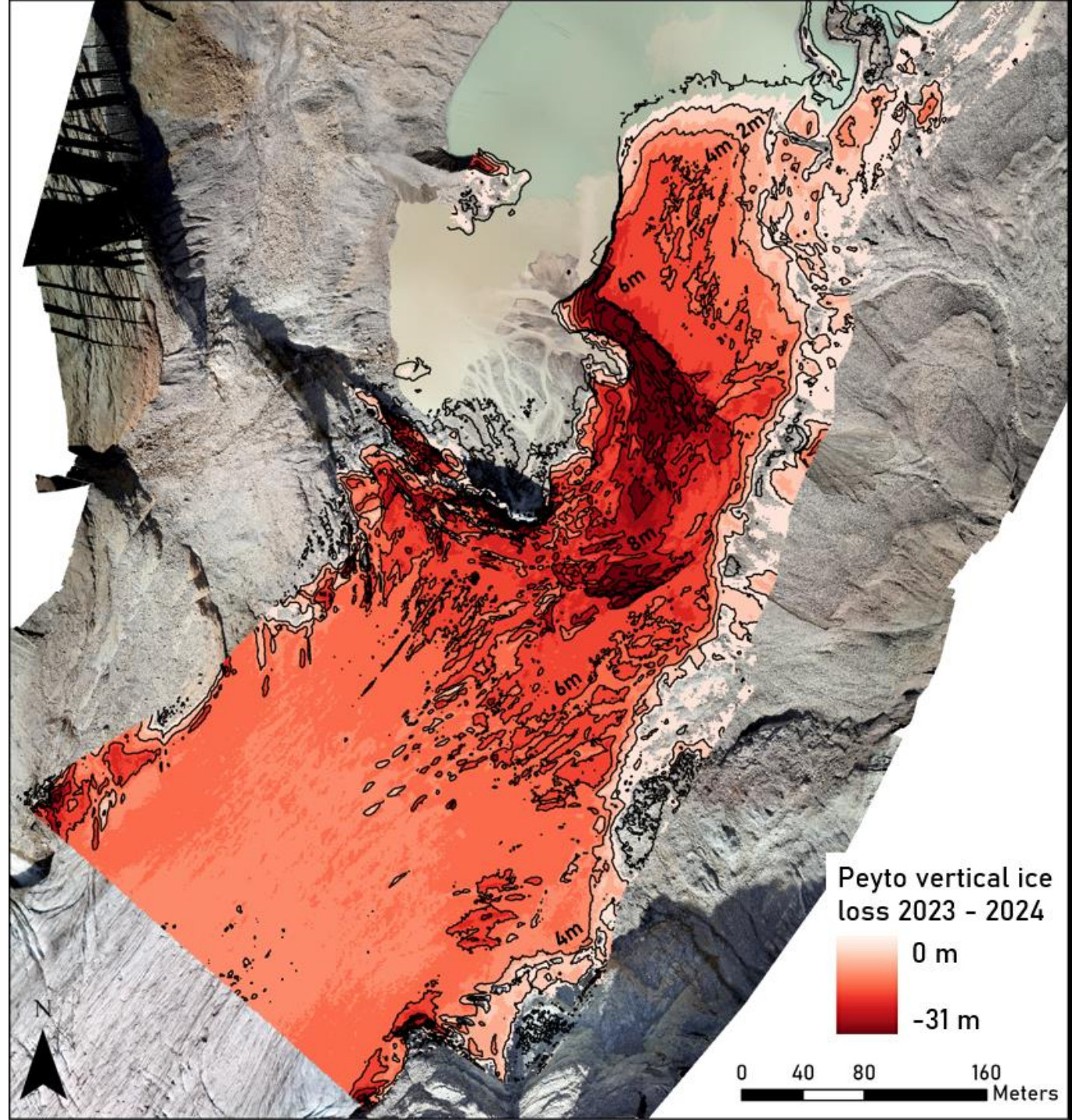
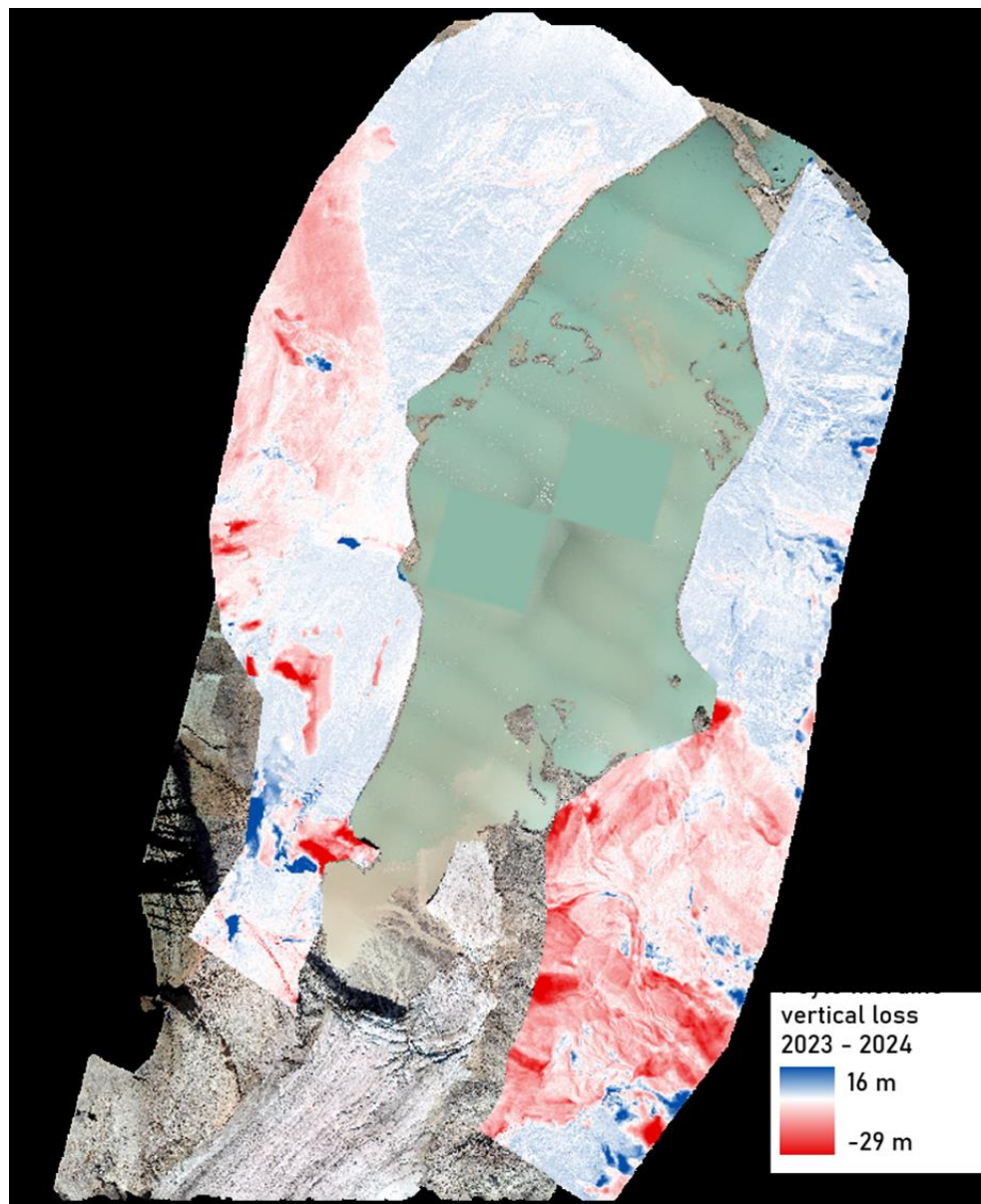
Extent of Peyto glacier toe from 2019 – 2024, overlain over aerial photography of Peyto taken in 2024

Definition of Peyto glacier toe for each year from 2019 - 2024

Peyto Glacier 5-year Vertical Loss



Peyto Glacier 1 Year Vertical Loss – ice and moraine

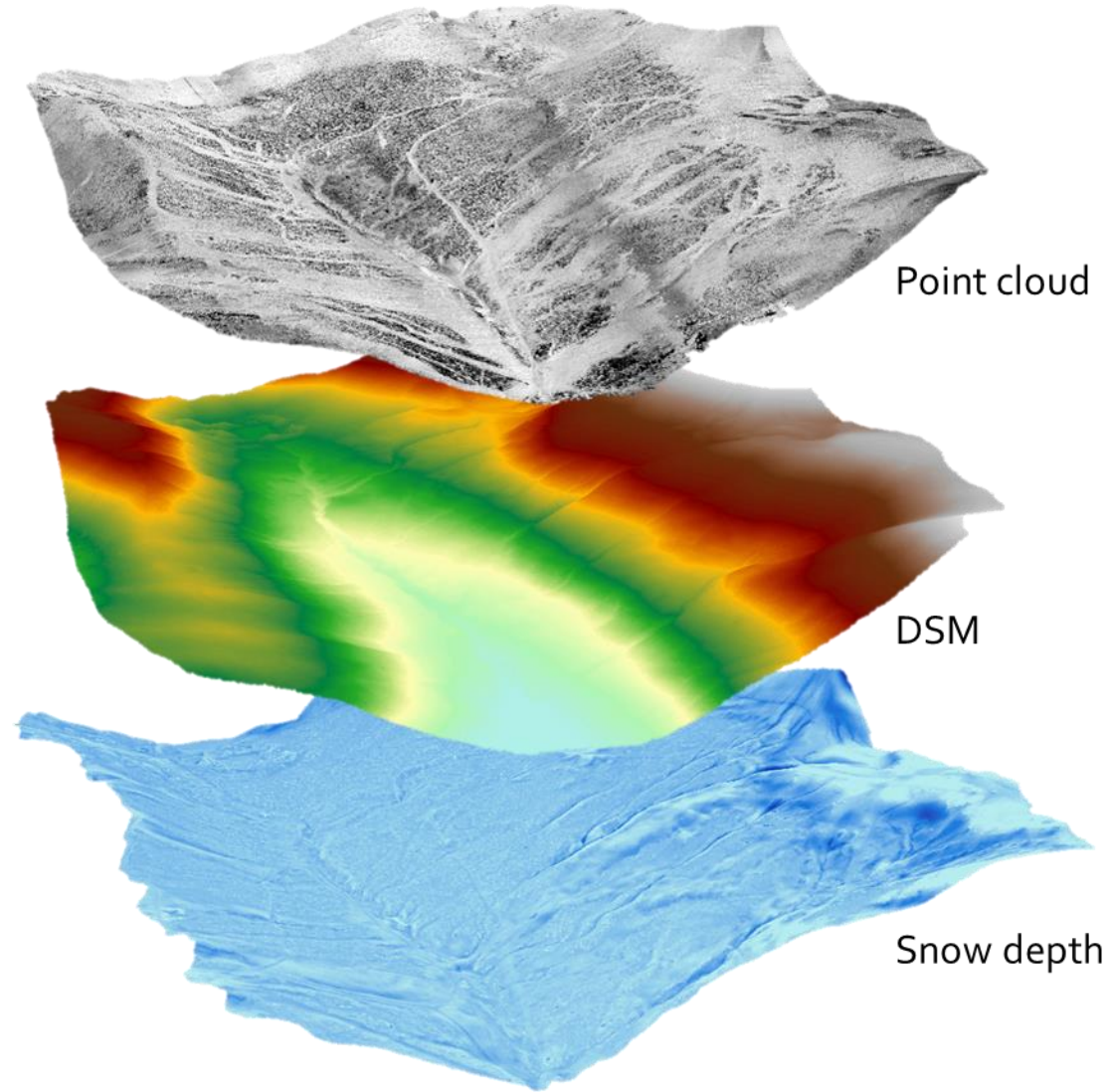
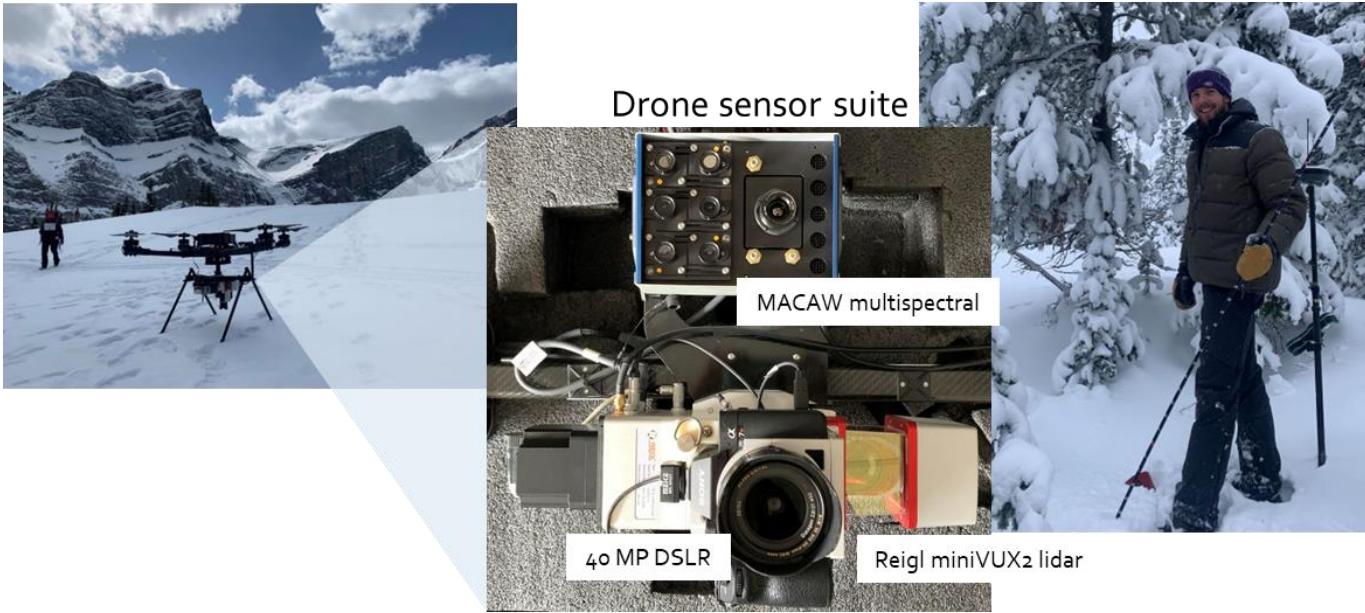


Peyto Glacier Change Summary

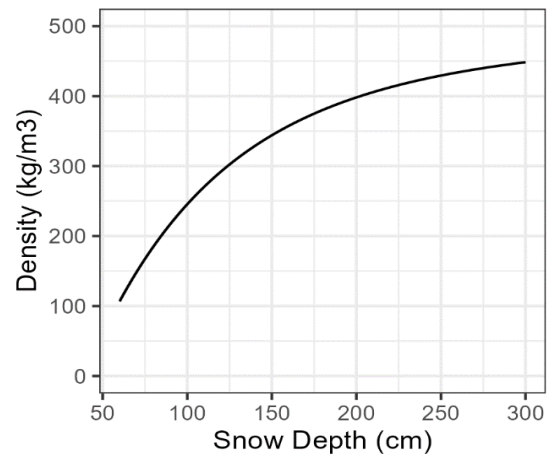
- Flow rates accelerated from 7.2 to 8.5 m/year over 2019 -2024
- 443 m of retreat since 2019
- 4 – 6.5 m of ablation due to melt in 2023-2024 – surface lowering of 31 m around ice surface collapse over internal conduit
- Substantial moraine surface change in 2023-2024 - from 29 m lowering to 16 m increase due to debris-covered ice melt and debris flows
- 30-35 m of ablation due to melt 2019-2024 with surface lowering of 56 m around ice surface collapse

$$\text{SWE (kg/m}^2\text{)} = \text{snow density (kg/m}^3\text{)} \times \text{snow depth (m)}$$

LiDAR provides snow depth, but snow density needed to calculate SWE

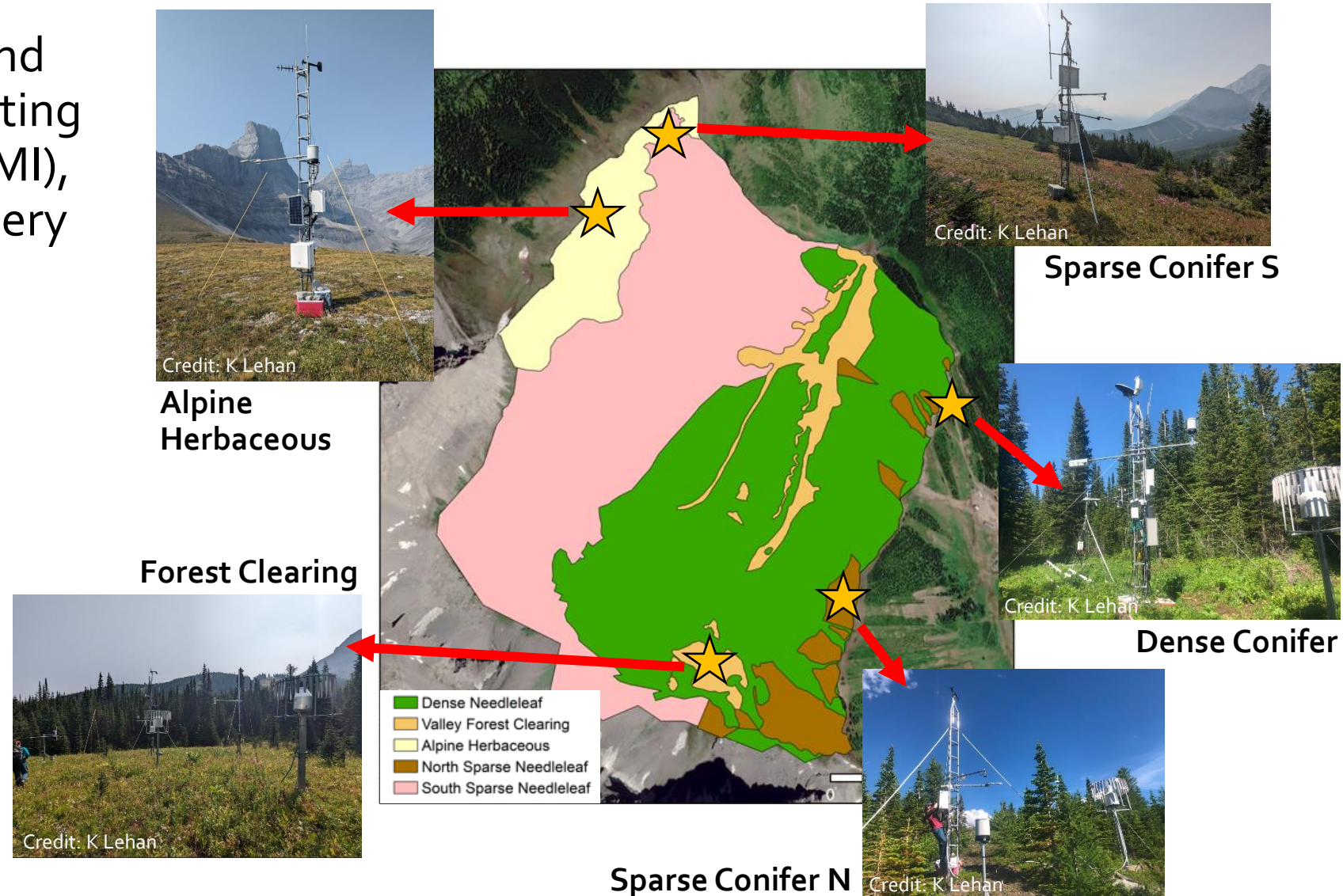


Pomeroy & Gray (1995) equation (winter)

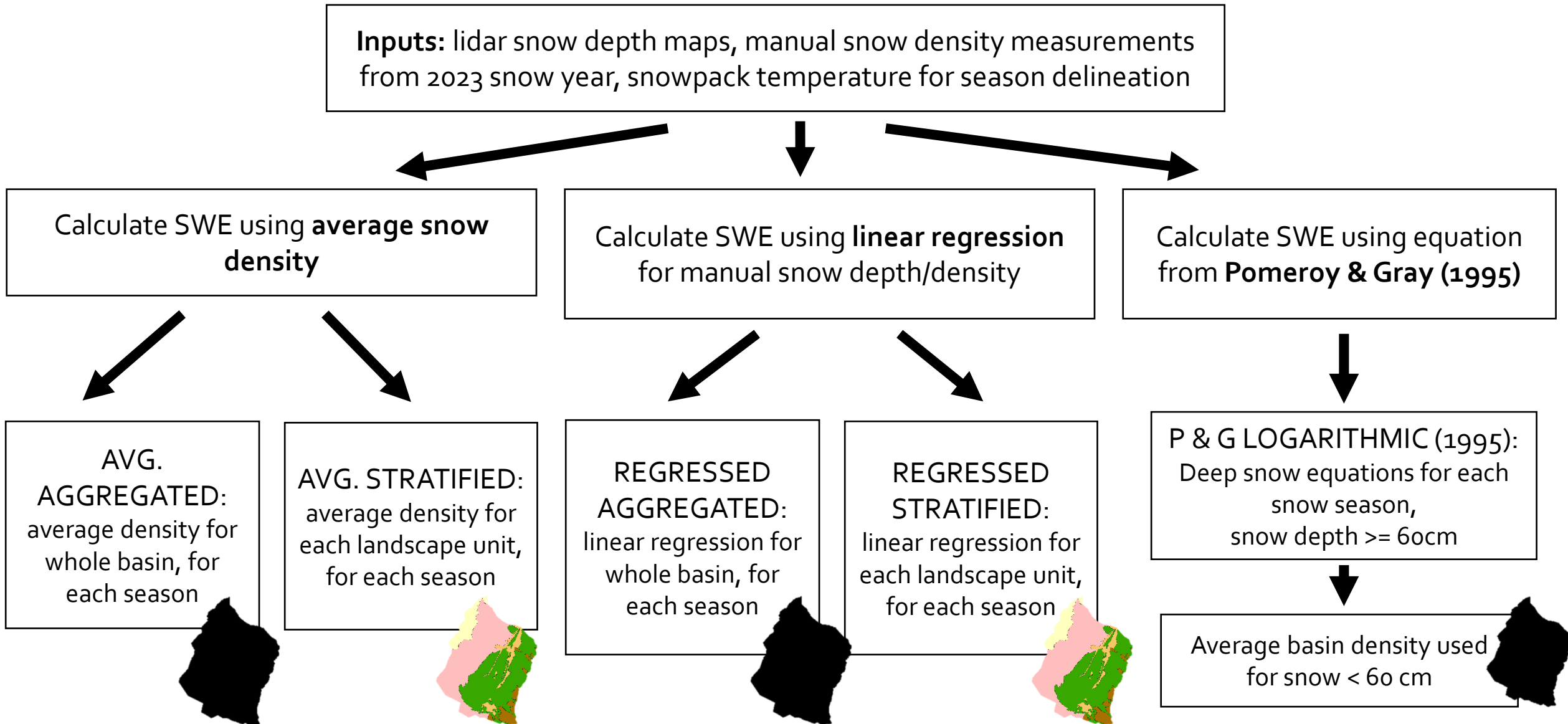


Landscape Units from Stratified SWE Sampling

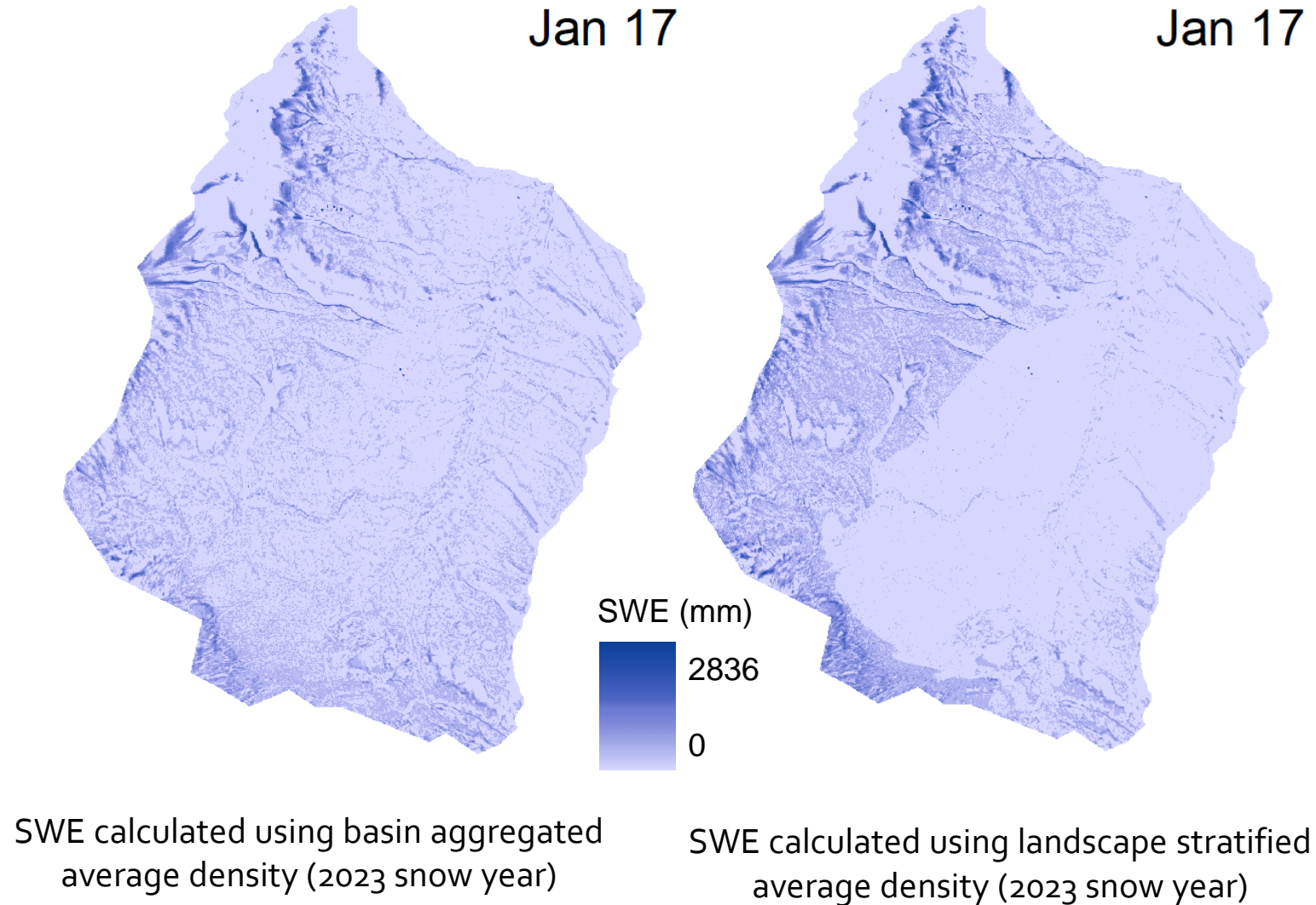
Landscape units (Steppuhn and Dyck, 1974) derived from existing landscape classifications (ABMI), and adjusted using RGB imagery and scientists' observations



SWE Calculation Workflow



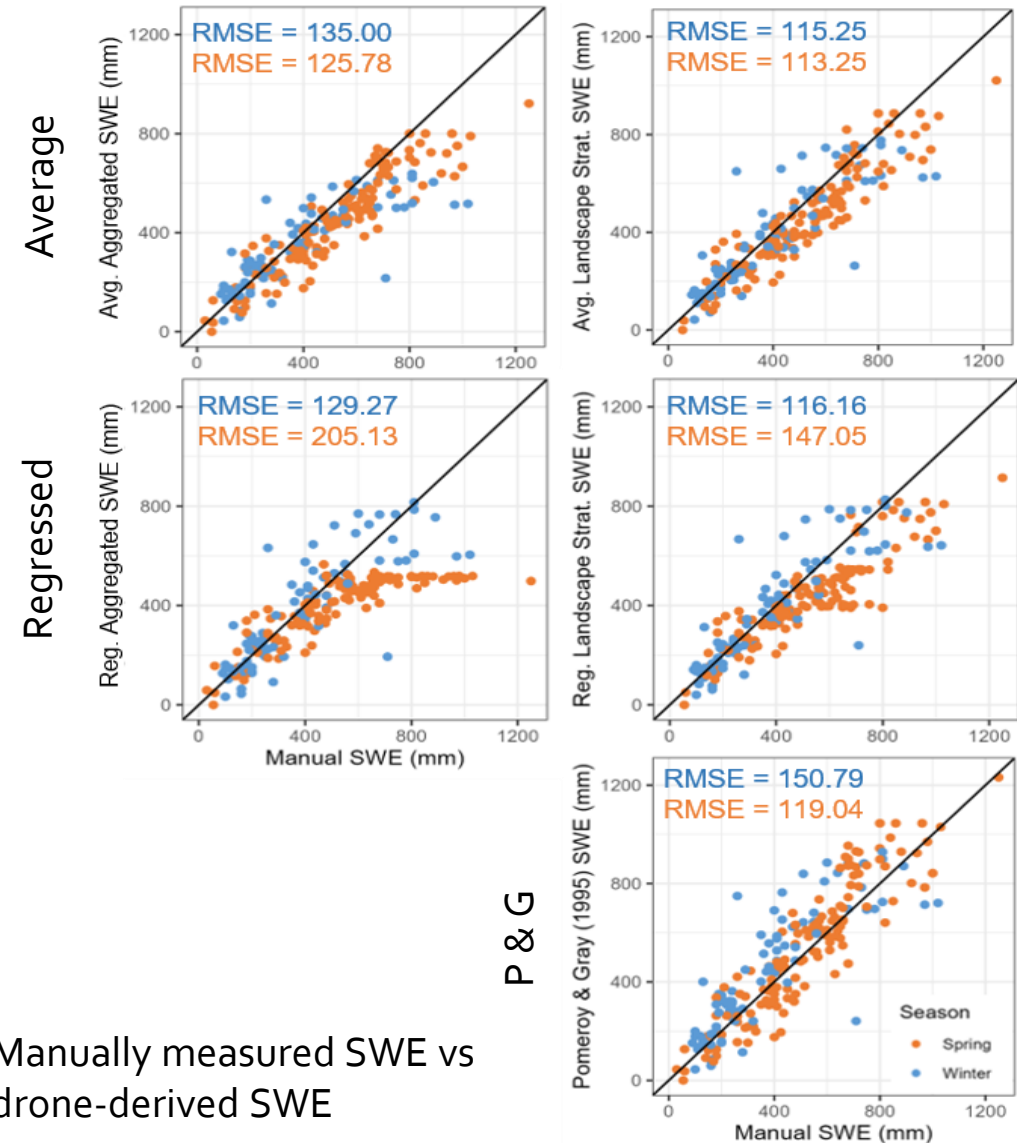
Methods – SWE Calculation Workflow



Results – SWE Point RMSE

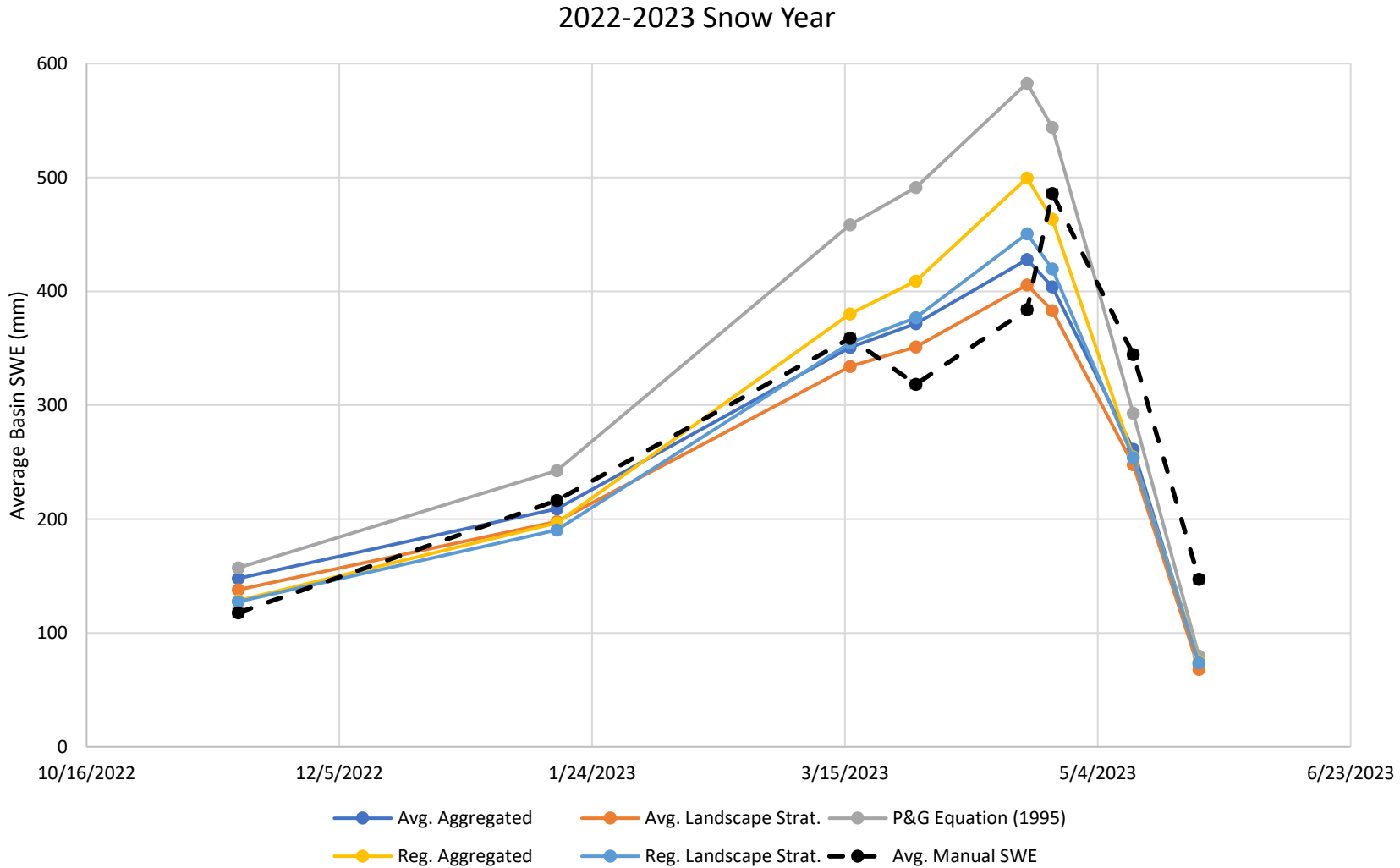
Plotting manual point SWE against calculated point SWE reveals:

- Regression methods fail in spring (expected!)
- Pomeroy and Gray (1995) equation surprisingly suitable in spring
- Landscape-stratified mean density provided the closest estimate of point SWE overall



Manually measured SWE vs drone-derived SWE

Areal-weighted Basin SWE



UAV-LiDAR snow survey summary

Applying 5 methods of density interpolation for basin-wide SWE calculation revealed:

Point SWE was best estimated by using a landscape stratified mean density (RMSE 114 mm, mean bias -51.9 mm)

- Landscape stratified methods outperformed aggregated approaches
- All methods failed to represent point SWE in the spring – P&G (1995) equation predicts spring point SWE best, but bias was large

Conclusions:

- i) **landscape-stratified snow density measurements improve SWE estimates from UAV-LiDAR snow depths**
- ii) **if density cannot be measured, Pomeroy and Gray (1995) equation can be used to estimate SWE in mountain environments**