TOWARDS DOWNSCALING OF PRECIPITATION PHASE FROM HIGH RESOLUTION METEOROLOGICAL FORECAST MODEL OUTPUT OVER COMPLEX



TERRAIN

Sabine RADANOVICS, Matthieu LAFAYSSE

Univ. Grenoble Alpes, Université de Toulouse, Météo-France, CNRS, CNRM, Centre d'Études de la Neige, Grenoble, France



Why to downscale precipitation phase?

- Plan to use the French NWP model AROME (Seity et al. 2011) and its ensemble forecast version PEAROME (Raynaud and Bouttier 2016) to provide surface meteorological variables in a future gridded snow cover forecast system at 250 m horizontal resolution.
- Model resolution:
- AROME 1.3km PEAROME 2.5km
- precipitation phase is crucial for snow cover modelling and highly dependent on altitude and atmospheric conditions

Precipitation phase at 250m resolution?

Height difference 2.5km vs. 250m resolution



• substantial altitude difference between pixels at different resolutions over complex terrain

Study area: Grandes Rousses Massif in the French Alps



Fig. 1: Left: France, center: zoom to the French Alps with pixel heights at the forecast model resolution of 2.5km, right: Zoom to the Grandes Rousses Massif with pixel heights of the 250m target grid

How to determine precipitation phase?

Precipitation type model diagnostic

• PEAROME model output contains a precipitation type classification based on the vertical wet-bulb temperature profile (Fig. 2)

Wet-bulb temperature iso-heights

Fig. 3: Height difference between the pixels at 250m resolution and the pixels at 2.5km resolution over the Grandes Rousses domain. Red: 250m pixel above 2.5km pixel, blue: 250m pixel below 2.5km pixel.



250m grid without taking height difference into account

Adjusted precipitation type using Tw iso-heights

- Snow melts at wet-bulb temperatures $(Tw) > 0^{\circ}C$
- Snow limit, i.e., altitude where snow is completely melted, is sometimes assumed at the height where $Tw = 1.5^{\circ}C$



For pixels with precipitation:

• dry snow and wet snow are kept if the pixel height is above the $0^{\circ}C$ Tw height

• melting snow is assigned if the pixel height is between the $0^{\circ}C$ Tw and the $1^{\circ}C$ Tw height

• mix of rain and snow is assigned if the pixel height is between the $1^{\circ}C$ Tw and the $1.5^{\circ}C$ Tw height

• rain if the pixel height is below the $1.5^{\circ}C$ Tw height



Fig. 5: Precipitation type from PEAROME forecast resampled at the 250m grid adjusted using Tw iso-heights

- using those decision rules mix-phase types occur in some valleys (Fig. 5)
- some Tw and height information included
- easy to apply, no extrapolation needed
- no quantitative information on solid and liquid precipitation fraction
- does not account for precipitation intensity

Fig. 2: Simplified Precipitation type classification from PEAROME forecast (3h lead time) over the French Alps for 1 April 2022 6h UTC. Wet-bulb temperature iso-lines: $0^{\circ}C$ in orange and $1.5^{\circ}C$ in red. Mix-phase types (melting snow and mix of rain and snow) mainly occur between those two Tw levels. However, drizzle also occurs at Tw $< 1.5^{\circ}C$, which illustrates that the snow level also depends on the precipitation mass to melt.

Latent heat method (Vionnet et. al (2022))

- Given a wet-bulb temperature profile, the available melting energy is calculated in each layer
- Snow limit is considered at the level where the accumulated energy is sufficient to melt all the precipitation falling through
- Accounts for precipitation intensity

Future plans and challenges

Apply the latent heat method for downscaling

Challenges:

- How to extrapolate the Tw profile below the original model pixel height in narrow valleys?
- Which assumptions on vertical temperature and humidity gradients in valleys are the most appropriate ones?

LATEX TikZposter

Contact

sabine.radanovics@meteo.fr

Centre d'Etudes de la Neige, Météo-France Domaine Universitaire, 1441 rue de la Piscine F-38400, Saint Martin d'Hères, France

References

• account for precipitation intensity

• possibly quantitative information on frac-

tion of solid and liquid precipitation

Advantages:

Raynaud, L., and F. Bouttier, 2016: Comparison of initial perturbation methods for ensemble prediction at convective scale. Quarterly Journal of the Royal Meteorological Society, 142 (695), 854–866, doi:10.1002/qj.2686. Seity, Y., P. Brousseau, S. Malardel, G. Hello, P. Bnard, F. Bouttier, C. Lac, and V. Masson, 2011: The AROME-France convective-scale operational model. Monthly Weather Review, 139 (3), 976–991, doi:10.1175/2010MWR3425.1. Vionnet, V., M. Verville, V. Fortin, M. Brugman, M. Abrahamowicz, F. Lemay, J. M. Thériault, M. Lafaysse, and J. A. Milbrandt, 2022: Snow level from post-processing of atmospheric model improves snowfall estimate and snowpack prediction in mountains. Under review